Miocene fire intensification linked to continuous aridification on the Tibetan Plateau

Yunfa Miao1,2,*, Fuli Wu2,*, Sophie Warny3, Xiaomin Fang2, Haijian Lu4, Bihong Fu5, Chunhui Song6, Xiaoli Yan6, Gilles Escarguel7, Yibo Yang8, Qingquan Meng6, and Piling Shi8

1Key Laboratory of Desert and Desertification, Northwest Institute of Eco-Environment and Resources, Chinese Academy of Sciences, Lanzhou 730000, China
2CAS Center for Excellence in Tibetan Plateau Earth Sciences and Key Laboratory of Continental Collision and Plateau Uplift, Institute of Tibetan Plateau Research, Chinese Academy of Sciences, Beijing 100101, China
3Department of Geology and Geophysics, and Museum of Natural Science, Louisiana State University, Baton Rouge, Louisiana 70803, USA
4Institute of Geology, Chinese Academy of Geological Sciences, Beijing 100037, China
5Institute of Remote Sensing and Digital Earth, Chinese Academy of Sciences, Beijing 100094, China
6School of Earth Sciences, and Key Laboratory of Western China’s Mineral Resources of Gansu Province, Lanzhou University, Lanzhou 730000, China
7Laboratoire d’Ecologie des Hydrosystèmes Naturels et Anthropisés, UMR 5023 LEHNA, CNRS, ENTPE, Université Lyon 1, 69622 Villeurbanne Cedex, France

ABSTRACT
Although fire is considered an important factor in global vegetation evolution and climate change, few high-resolution Miocene fire records have been obtained worldwide. Here, two independent micro-charcoal-based fire records from the northern Tibetan Plateau were analyzed; both show similar trends in micro-charcoal concentrations through time, with low abundances in the warmer Middle Miocene Climate Optimum (18–14 Ma) followed by a continuous increase throughout the late Miocene (14–5 Ma) cooling. Our detailed statistical analyses show that the micro-charcoal concentration trend is highly positively correlated to the trend in oxygen isotopes (δ18O, r = 0.94) and xerophytic species (%xer, r = 0.95). We propose that this indicates a link between fire frequency and climate. The charcoal resulting when plant matter is burned has been referred to as “fossils of fire” and are regarded as one of the most important proxies in the study of fire history (e.g., intensity, frequency, etc.) (Patterson et al., 1987; Whitlock and Larsen, 2002; Bowman et al., 2009; Bond, 2015). Here, we analyze fire histories based on relatively high-resolution micro-charcoal records from two independent sites on the northern Tibetan Plateau: the KC-1 core (38°03′N, 91°45′E; drilled by the Qinghai Petroleum Administration Bureau) in the Qaidam Basin, characteristic of a lacustrine environment (Miao et al., 2011), and the KM/KNX sections in the Kumkol Basin, characteristic of a fluvial-lacustrine environment (Lu et al., 2016, 2018) (Fig. 1).

STUDY SITES AND DATA
Situated in southern Inner Asia, the northern Tibetan Plateau is marked by a typical inland hyperarid climate, with mean annual precipitation of <200 mm. As a result of its high elevation, the area is also very cold and windy. The Westerns represent the dominant atmospheric circulation pattern throughout the year, while the East Asian summer monsoon generally reaches only the southeastern Qaidam Basin (Wu et al., 1985) (Fig. 1). Under these conditions, bare ridges and desert vegetation constitute the typical land cover. The desert vegetation that grows on gravelly lake margins and on well-drained soils of alluvial fans is composed mainly of Ephedra, Nitraria, Tamarix, Poaceae, Cyperaceae (as well as Asteraceae), and occasionally Picea and Sabina on the high mountains near the eastern part of the Qaidam Basin.

Tectonically, the ~1600-km-long Altyn Tagh fault separates the low-lying Tarim Basin to the north from the Tibetan Plateau. The left-lateral strike-slip Kunlun fault marks the northern boundary of the high-elevation, low-relief main part of the Tibetan Plateau, extending for ~1500 km. A broad triangular area consisting of the Kumkol Basin, the Qimian Tagh range, and the Qaidam Basin is bounded to the northwest by the Altyn Tagh fault and to the south by the Kunlun fault (Fig. 1). Previous work based on apatite fission-track modeling has shown that the Qaidam and Kumkol Basins have been separated by the synchronous initial uplift of Qimian Tagh at ca. 40–30 Ma (Liu et al., 2017).

The present Qaidam Basin, with an average elevation of ~3000 m, is the largest intermontane basin on the northeastern Tibetan Plateau, and

INTRODUCTION
Late Cenozoic global fire activity has been considered as one of the key factors influencing modern vegetation evolution and climate change (Bowman et al., 2009; Dale et al., 2001; Edwards et al., 2010). However, globally, few Miocene high-resolution records have been obtained with which to assess fire occurrence (Hoetzel et al., 2013). As part of the second-largest arid to semi-arid area in the world, the northern Tibetan Plateau is a unique location for studying fire history, along with vegetation and aridity evolution (Miao et al., 2016); however, the region’s numbers of fire records, sampling resolution, combustion source data, and controlling factors still need further investigation.

The charcoals resulting when plant matter is burned have been referred to as “fossils of fire” and are regarded as one of the most important proxies in the study of fire history (e.g., intensity, frequency, etc.) (Patterson et al., 1987; Whitlock and Larsen, 2002; Bowman et al., 2009; Bond, 2015). Here, we analyze fire histories based on relatively high-resolution micro-charcoal records from two independent sites on the northern Tibetan Plateau: the KC-1 core (38°03′N, 91°45′E; drilled by the Qinghai Petroleum Administration Bureau) in the Qaidam Basin, characteristic of a lacustrine environment (Miao et al., 2011), and the KM/KNX sections in the Kumkol Basin, characteristic of a fluvial-lacustrine environment (Lu et al., 2016, 2018) (Fig. 1).

covers an area of ~120,000 km². The sediments recovered from the KC-1 core, reaching 3425 m in depth, were extracted from the western basin and span ages of 18 Ma to 5 Ma, with an average sedimentation rate of ~260 m/m.y. (Miao et al., 2011) (Fig. 2A). Another section, named NG (Naoge), from the eastern basin and spanning ages of ca. 18–10 Ma, was also used for comparison (for details, see Miao et al., 2016). The rhombus-shaped Kumkol Basin has an average elevation of ~4200 m and an area of ~17,500 km² (Fig. 1). We sampled the KM section, spanning 2869 m in thickness, and the KNX section, spanning 1125 m, from the central basin. Chronostratigraphic analysis of these two sections yielded ages of 17.5–6.3 Ma in the KM section (Lu et al., 2016), and 8.2–4.2 Ma in the KNX section (Lu et al., 2018), providing average sedimentation rates of ~260 m/m.y. and ~290 m/m.y., respectively (Fig. 2B). The composite age range of 17.5–5.0 Ma for these two sections (herein KM/KNX) was used for the micro-charcoal analyses and comparison with the KC-1 core.

**METHODS AND RESULTS**

In this study, 188 samples from the KC-1 core (Qaidam Basin) and 250 samples from the KM/KNX sections (Kumkol Basin) were used for micro-charcoal identification and processed via standard palynological techniques. A known number of Lycopodium clavatum spores were initially added to each sample for calculating the micro-charcoal concentration (MC) (Miao et al., 2016). Two basic micro-charcoal shapes (sub-long [L] and sub-round [R]) were identified by calculating the ratio of length (major axis) to width (minor axis); if the value was >2.5, the micro-charcoal grain was classified as L, and if <2.5, as R. The MCs for both types (MC<sub>L</sub> and MC<sub>R</sub>, respectively) were determined, and then added together to obtain the total MC values, referred to as MC<sub>total</sub> (Table DR4 in the GSA Data Repository<sup>1</sup>). The MC<sub>total</sub> for the previously studied NG section in the eastern Qaidam Basin (Miao et al., 2016; Fig. 1) was also included in the statistical analyses.

First, the existence of a trend in the three separate time series (KC-1, KM/KNX, and NG) of observed MC<sub>total</sub> values was tested using the Mann-Kendall trend test (Table DR1). Then Pearson correlations were computed for the three possible pairs of MC<sub>total</sub> time series based on spline-fitted log-transformed MC<sub>total</sub> values interpolated every 50 k.y. along each time series, resulting in three interpolated data sets described within the same time frame (Table DR2). Last, a composite MC<sub>total</sub> time series was compiled from the three separate series, and finally correlated to the time series of marine δ<sup>18</sup>O, atmospheric pCO<sub>2</sub>, and percentages of xerophytic taxa in the Qaidam Basin using the same interpolation procedure based on spline-fitted curves (Table DR3). In all cases, correlation significance was evaluated through a Mantel test based either on the interpolated MC<sub>total</sub> values or on the time autocorrelation–free first-order differences between “raw” interpolated MC<sub>total</sub> values. All computations were done using PAST software, version 3.16 (https://folk.uio.no/ohammer/past/; Hammer et al., 2001).

The most notable characteristics between the KC-1 and KM/KNX sites are their similar patterns (Pearson r = 0.95; Table DR2) but different absolute values of MC<sub>L</sub>, MC<sub>R</sub>, and MC<sub>total</sub>. The KC-1 core has the highest grain counts, varying between 46,300 and 900 grains/g, with an average of ~7000 grains/g (Fig. 2A). The KM/KNX sections vary between 36,100 and 200 grains/g, with an average of ~4000 grains/g (Fig. 2B). Overall, the trends of MC<sub>total</sub> in the three sections (including NG) are all highly significantly increasing with time (Figs. 3B–3D; Tables DR1 and DR2).

As a result, the composite MC<sub>total</sub> time series also shows a highly significant trend (Mann-Kendall...
Figure 3. Data time series of climatic proxies and tectonic events on the northern Tibetan Plateau. MMCO—Middle Miocene Climate Optimum. A: Global deep-sea δ18O records (Zachos et al., 2001) (light blue) with 12-point averaging (darker blue). B–D: Micro-charcoal concentration (MC) records from KC-1 core, western Qaidam Basin (this study) (B), KM/KNX sections, Kumkol Basin (this study) (C), and NG section, eastern Qaidam Basin (Miao et al., 2016) (D). E: Xerophytic pollen percentages in KC-1 core (Miao et al., 2011). F: Synthesis of published pCO2 proxy data. Boron (diamonds) from Foster et al., 2012; Greenop et al., 2014. Stomata (inverted triangles) from Küppers et al., 2008. Alkenone (regular triangles) from Zhang et al., 2013. B/Ca (squares) from Tripati et al., 2009. Paleosol (circles) from Ji et al., 2018. Error bars represent uncertainties in underlying assumptions of each proxy. G–H: Ratios of MC/NC (L—sub-long grains; R—sub-round grains) from KC-1 core and KM/KNX sections, respectively (this study). I: Tectonic events on the northern Tibetan Plateau (after Miao et al. [2012], Li et al. [2014], and Chang et al. [2015]). Curves in B–D and G–H are by three-point averaging.

**DISCUSSION**

**Fire and Aridification**

Based on the relationship between MC_{total} and fire (Hoetzl et al., 2013; Patterson et al., 1987; Bond, 2015), higher values of MC_{total} over long time scales (e.g., millions of years) can be related to increased fire activity, e.g., greater frequency or intensity. The MC_{total} data from the KC-1 core and KM/KNX sections (Figs. 3B and 3C), as well as from the NG section (Miao et al., 2016) (Fig. 3D), support the view that fire activity on the northern Tibetan Plateau steadily increased during the middle to late Miocene. The overall average expansion of xerophytic vegetation in the western Qaidam Basin (Miao et al., 2011), as well as an increase in MC_{total}, support an increase in fire activity and suggest that these two phenomena are strongly linked from 18 to 5 Ma (Miao et al., 2016). A relatively low fire frequency and/or strength occurred during the relatively wet Middle Miocene Climate Optimum (MMCO, 18–14 Ma), while an overall increasing trend is observed during the drier late Miocene (14–5 Ma).

**Fire Distribution**

Although the relationship between fire intensification and aridification on the northern Tibetan Plateau correlates well, the question of fire distribution remains. It is very important to understand whether the fire intensification is due to an enlarged combustion area or an increase of combustion frequencies.

Studies show that the L and R in micro-charcoal shapes represent grass and wood sources, respectively (Umbanhowar and McGrath, 1998; Daniau et al., 2013; Crawford and Belcher, 2014). This means that the samples with higher MC_L seem to be dominated by the combustion of grasses, while the samples with higher MC_R seem to be dominated by combustion of wood. Therefore, the ratios of MC/LMC_R can be directly used to indicate the grass/wood ratios during combustion, with higher values indicating more grass than wood, and vice versa. In this study, the higher values of the MC/LMC_R in the KM/KNX sections (~0.6–1.0) are related to combustion dominated by grass (Fig. 3H), while lower values in the KC-1 core (~0.5) suggest relatively higher wood combustion (Fig. 3G). Interestingly, although the MC/LMC_R values are different, the trends remained stable over the examined interval for both sections. Therefore, despite differences in vegetation source material between the two sections, and the expansion of grasslands during late Miocene cooling, the forest-steppe ecotone must have still been the dominant environment where fire occurred in order for such stable trends in MC/LMC_R values to have been maintained. The conceptual renderings in Figure 4 illustrate this phenomenon. During the MMCO, the area of the forest-steppe ecotone was large but fire frequency was low (Fig. 4A) under the relatively wet climate (Miao et al., 2011), and less micro-charcoals were produced (Fig. 4A). In contrast, during the late Miocene, the forest-steppe ecotone remained a fire source area but likely experienced an overall decreasing trend in areal extent as xerophytic vegetation expanded under a drying climate (Miao et al., 2011). However, the persistently increasing fire frequency due to the drier conditions produced more charcoal particles, thereby greatly counteracting the negative effect of eco-tone shrinkage (Fig. 4B).

**Global Cooling, CO2, and Tectonics**

After compiling global low-sampling-resolution fire records, Bond (2015) argued that none of the potential global factors (oxygen, rainfall seasonality, CO2, novel flammable growth forms) could adequately explain the spread of increased fire activity over long Cenozoic time scales. In this study, the increasing fire frequency is strongly negatively correlated to Bond’s compiled results of atmospheric pCO2 (r = -0.86; Fig. 3F) and is positively closely linked to the global oxygen isotope record (Zachos et al., 2001) (r = 0.94; Fig. 3A) and the expansion of xerophytic vegetation (r = 0.95; Miao et al., 2011; Fig. 3E).

CO2 has been argued as having an important role in fertilization of vegetation through affecting photosynthetic assimilation (Wolfe and Erickson, 1993), thus changing the amount of annual gross primary production (Sun et al., 2014). Theoretically, high pCO2 facilitates
CONCLUSION

A series of detailed micro-charcoal data sets obtained from two new independent sites on the northern Tibetan Plateau (from the KC-1 core in the western Qaidam Basin) and previously published data from the NG section in the eastern Qaidam Basin, offer a perspective on the fire history during 18–5 Ma on the northern Tibetan Plateau. Our results show increased fire occurrence based on low values of micro-charcoal during the warm and dry-late MMCO to high values during the cooler and drier late Miocene. Increased fire frequencies in the forest-steppe ecotone counterbalanced the concurrent reduction in combustion area. Our data and statistical analyses support the view that fire activity on the northern Tibetan Plateau is primarily linked to global cooling, with decreased atmospheric pCO₂ reflecting the continuous aridification on the northern Tibetan Plateau at 18–5 Ma. The influence of tectonic activity on the northern Tibetan Plateau is likely a secondary controlling factor.

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